

planning efforts and may be used as the framework to identify regional eolian processes and to guide future studies. For additional discussion regarding the Imperial Dunes area and other mapped areas, please refer to the project report

(Lancaster and Bedrossian, 2014).

Imperial Dunes Area

The Imperial Dunes area is located entirely within the Colorado Desert geomorphic province, in Imperial County. The Colorado Desert geomorphic province is a low-lying barren desert basin dominated by the Salton Sea. The province is a depressed block between active branches of alluvium-covered San Andreas Fault with the southern extension of the Mojave Desert on the east. The western portion of the map area is characterized by the ancient beach lines and silt deposits of extinct Lake Cahuilla (CGS, 2002). The climate is arid, hot, and dry with average annual temperatures reported for Calexico ranging between 13.3° and 30.1°C (55.9° and 86.2° F) (DRI, 2014). According to the Desert Research Institute (DRI), January mean temperatures range from 3.9° to 19.6°C (39.0° to 67.2°F) and July mean temperatures range from 6.6° to 39.9°C (75.8° and 103.9°F). The average annual rainfall is 68.3 mm/yr (DRI, 2014).

The northwest trending Algodones Dunes form the prominent feature in the map area. As one of the largest active dune fields in the United States, they are up to 11 km wide and roughly 72 km long and extend from Mammoth Wash in the northwest into the northernmost part of Baja California. The Algodones dune field rests on the margin of an ancient lake bed that forms a topographically planar area called the East Mesa. The East Mesa area also contains prominent linear dunes, coppice dunes, and sand sheets. Both the Algodones and East Mesa dunes are dominated by the northwest-trending, fault-bound Salton Trough, which includes the Salton Sea to the northwest and the Imperial Valley, a broad 48 km wide, sparsely populated alluvium-filled valley that lies largely below sea-level to the south of the Salton Sea (Bedrossian, 2011). Below the valley alluvium lie 6100 m of Tertiary and Quaternary sedimentary rocks, predominantly non-marine sandstones, clays, and lacustrine deposits. This thick sequence of sediments is thought to have been derived from the Colorado River during several episodes of prehistoric deposition, i.e., whenever the river shifted its course away from the Gulf of California, where it now flows, to the Salton Trough (Muhs et al., 1995). Numerous active faults within and bordering the Salton Trough suggest it is a depressed block caused by seafloor spreading that is still widening the Gulf of California to the south. South of the Salton Sea, within the Imperial Valley, the western and eastern boundaries of the trough are characterized by ancient beach lines and silt deposits of the now extinct Lake Cahuilla. Late Pleistocene shorelines are thought to underlie the Algodones dune field on East Mesa and the late Holocene lacustrine shorelines of Lake Cahuilla are generally found at elevations of about +12-14 meters above sea level (Muhs et al., 1995). The East Mesa and Algodones dune field are bordered on the east by the Chocolate and Cargo Muchacho mountains. These mountains are composed of Precambrian igneous and metamorphic rocks, Mesozoic granitic and metamorphic rocks, and Tertiary volcanic rocks (Jennings, 1967; Muhs et al., 1995). Holocene alluvial wash and alluvial fan deposits, consisting mostly of gravel and sand, comprise the broad gently sloping piedmont between the mountains to the east and the East Mesa area.

Dune forms in the Algodones Dunes have been well described in studies by numerous authors including Norris and Norris (1961), Olmsted et al. (1973), Sharp (1979), Smith (1982), Kocurek and Nielson (1986, Nielson and Kocurek (1986), Havholm and Kocurek (1988), Sweet et al. (1988) and Dohrenwend and Smith (1991). Kocurek and Nielson (1986) mapped an east-west transect across the northwestern end of the Algodones Dunes and described the following sequence of dune forms: (1) sand sheet and zibars; (2) linear dunes, with zibars in interdune corridors; (3) compound-complex crescentric ridges or draa (farther to the southeast); and (4) an eastern sand sheet. CGS's current mapping is in general agreement with these four zones. In aerial photographs and on GoogleEarth imagery, coppice dunes, along the eastern and western margins of the sand sheets, and linear dunes and transverse or crescentric ridges are also present throughout the dune field (Muhs et al., 1995). Linear dunes and sand streaks (McKee, 1979) are present and visible on the East Mesa.

Interpretation and Description of Map Units

In this map area, CGS focused preliminary mapping efforts on the identification of Holocene age eolian deposits. Mapping included the identification of features related to primary geomorphic processes (e.g. alluvial fan), followed by interpretation of those deposits related to secondary processes (e.g. eolian sand sheets superimposed on alluvium). CGS made field observations and collected data in March and April 2014. These observations included identification of soil texture and pedogenic development, as well as descriptions of surface cover, including the presence of gravel lag, desert pavement, or the accumulation of freshly deposited eolian sands. Based on field observations, Quaternary alluvial deposits exhibiting soil development and/or having a gravel lag were not retained as map units. These are Pleistocene to mid-Holocene age deposits that have been isolated from deposition for thousands of years. Eolian transport may occur across these surfaces, but thick, laterally continuous deposition of eolian sand is of low potential.

The collection of site-specific wind data was not a part of the project scope, however, CGS used a relatively robust collection of published wind data in analyzing sand transport for the area. Utilizing sand rose data from four weather stations in the vicinity of the Algodones Dunes and one locality at the junction of the All-American and Coachella Canals, immediately west of the southern end of the Algodones Dunes (Sweet et al., 1988), Muhs et al. (1995) determined that: winds form a low energy regime in Indio and are from the northwest all year; winds form an intermediate energy wind regime at El Centro, in which winds are from the west most of the year and sand drift is predominantly west to east; at the junction of Coachella and All-American Canals, the overall drift of sand was from the northwest to southeast. Based on the dune features mapped by Kocurek and Nielson (1986), Muhs et al. (1995) mapped the orientations of linear dune crests, barchanoid ridge (draa) brinks (top of slip face or crest), and transverse ridge (draa) brinks in the Algodones Dunes, and crests of linear dunes and streaks on East Mesa, with the following results:

- Active linear dunes (linear sief) in the Algodones Dunes have a mean orientation of N43°W±12° (n=89), which approximates the dominant dune-forming wind direction. Transverse and barcanoid ridges in the Algodones Dunes have a mean brink orientation of N45° $E\pm16^{\circ}$ (n=57), also implying a dominant dune-forming wind direction from the northwest, about N45°W. Linear dunes and streaks on East Mesa have a mean orientation of N71°W \pm 9° (n=255), although the southernmost dunes on East Mesa have an orientation that is closer to the mean orientation of the Algodones
- Because the East Mesa dunes appear to have been stabilized only within the past few hundred years or less, the differences in resultant drift directions on East Mesa may be due to significant differences in resultant drift directions in different parts of the Salton Trough, even over short distances.

To evaluate possible eolian sand sources for the Algodones Dunes and East Mesa, Muhs et al. (1995) performed simple mineralogical studies on a total of 114 samples. Samples were taken from a variety of dune types, sand sheets, alluvial fans, floodplains, and river channels. Semi-quantitative estimates of the relative amounts of quartz, potassium (K)-feldspar, and plagioclase in Algodones and East Mesa eolian sediments were made using x-ray diffractometry (Muhs et al., 1995). Results were then compared with data on the composition of Colorado River sediments previously studied by Van Andel (1964), Olmstead et al. (1973), and Van de Kamp (1973). These studies showed that Colorado River sediments are characterized mostly by quartz (65 to 75%), with lesser amounts of feldspar (10 to 20%), and rock fragments (1 to 8%). Data from Muhs et al. (1995) indicated that Lake Cahuilla sediments have a similar composition to the Colorado River deposits, but contain significantly more quartz and less plagioclase than San Bernardino Mountains-derived alluvium. Muhs et al. also found that Colorado River and Lake Cahuilla sediments are generally higher in quartz than alluvial deposits derived from the Chocolate Mountains. Although their findings differed somewhat from Van de Kamp (1973) who found lower quartz abundances in the Lake Cahuilla sediments than in the Colorado River sediments, Muhs et al. (1995) concluded the Lake Cahuilla sediments were most likely derived from Colorado River, with possibly some contributions from the Chocolate Mountains, but with little or no contribution from the San Bernardino Mountains. Their data also indicated that, since there were no significant differences between eolian sediments in the Algodones Dunes and East Mesa, the two areas appear to have a common origin, i.e., eolian sediments in the Algodones Dunes and in East Mesa were derived from Lake Cahuilla sediments, which in turn were derived mainly from the Colorado River. Based on their wind, mineralogical, and trace element studies, combined with data provided by Havholm and Kocurek (1988), Muhs et al. (1995) concluded that (1) the dominant dune-forming winds are from the northwest on both the Algodones and East Mesa dune fields, (2), that the Lake Cahuilla shoreline deposits both west and northwest of the dunes are the immediate source of the eolian sediments; and (3) that the northwest prevailing winds during fall, winter, and spring are responsible for the movement of eolian sand from the shore line to the dune fields.

Field observations by CGS indicated that while eolian transport occurs across the East Mesa area, between the former Lake Cahuilla deposits and the Algodones dune field, there are local differences in accumulation activity. The northern and central region of East Mesa was perhaps actively accumulating eolian deposits as linear dunes, coppice dunes, and sheets after the desiccation of Lake Cahuilla, about 300 years before present (Waters, 1983). However, field observations indicate, that in part, dunes and sand sheets in this area are stabilized by vegetation. These eolian deposits exhibited no soil development (map unit Qe2 and Qe2/Qal) and show locally active accumulations of eolian sand, wind abrasion (erosion), entrainment, and transport. Because the East Mesa dunes appear to have been stabilized only within the past few hundred years or less, Muhs et al. (1995) explored the use of an index of dune mobility (M), originally developed by Lancaster (1988). Using 11 localities in or adjacent to sand dunes or sand sheets in the Colorado Desert, Muhs et al. (1995) found the region should have fully active dunes. The only exception to the predicted and observed degree of activity was the stabilized eolian deposits on East Mesa. Based on various water table elevation studies reported by Loeltz et al. (1975), and the review of aerial photographs, Muhs et al. concluded that stabilization of the East Mesa Dunes and sheets may have been caused by recent rises in the water table because of local recharge of ground water from irrigation and canal leakages since 1939. An additional explanation could be that the source of eolian sand in this area has been reduced given the change in land use over the last century to west of the East Mesa area. Conversely, the southern portion of East Mesa has a laterally extensive body of active sand sheets and dunes. CGS reviewed both the most recent USDA soil maps (1981) as well as soil data developed for the Imperial area in 1903. In general, the eolian deposits on the East Mesa align with soils mapped in the Rositas Series (Soil Survey Staff, 1981). Mapping in 1903 of the lowlands to the west of the mesa, indicates that the valley's surficial deposits are comprised of Dune Sand, Imperial Sandy Loam, Imperial Sandy Loam, Imperial Loam and Imperial Clay. Of these, the sand and sandy loam soils are described as being of eolian source (Soil Survey Staff, 1903). The general locations of these soils have been digitized, grouped, and presented on the map as a cross hatched pattern. While they lie within an area of agricultural land use that is mapped as disturbed (D), they give an indication of potential eolian sources prior to the development of the area for agricultural purposes.

The Algodones dune field is an effective barrier to sediment transport from alluvial wash and alluvial fans onto the East Mesa area. Sediment is impounded on the northern margin of the dune field from Mammoth Wash to the Mexican border. Alluvial embayments occur along the northern margin of the dune field, suggesting that sediment from the Chocolate and Cargo Muchacho mountains is incorporated into the dunes along the northern margin. This observation is consistent with the conclusion by Muhs et al. (1995) that there is some contribution from the Chocolate Mountains. Sand sheets and coppice dunes mantle the surface within the embayments along the northern margin (Map unit Qe3/Qal). Mammoth Wash forms the western boundary of the Algodones Dunes. This wash and many others to the west are incised into older alluvium (Qol and Qoa), becoming distributary alluvial fans (Qf) that prograde onto the younger lake bed deposits of Lake Cahuilla (Ql). Sand transport occurs from the source Lake Cahuilla deposits (Ql) across the surface of the older alluvium and into the western tip of the dune field. Eolian sand accumulations begin to thicken a mile west of Mammoth Wash (Qe3/Qal), transitioning into the thick deposits with dune form to the east of Mammoth Wash (Qe3). The following description of map units summarize CGS' observations.



ccurs between mapped source and active eolian deposits

REFERENCES

Bedrossian, T. L., Roffers, P., Hayhurst, C. A., Lancaster, J. T., and Short, W. R., 2012, Geologic Compilation of Quaternary Surficial Deposits in Southern California, California Geological Survey, Special Report 217, 20 p, 25 plates, scale 1:100.000 Blair, T. C., and McPherson, J. G., 1994, Alluvial Fans and Their Natural Distinction From Rivers Based on Morphology, Hydraulic Processes, Sedimentary Processes, and Facies Assemblages: Journal of Sedimentary Research, 64(3). Bryant, W. A. (compiler), 2005, Digital Database of Quaternary and Younger Faults from the Fault Activity Map of California, version 2.0. California Geological Survey (CGS), 2002, Note 36: California Geomorphic Provinces, California Department of Conservation, California Geological Survey, 4p. Cooke, R. U., Warren, A., and Goudie, A. S, 1993, Desert Geomorphology: University College of London Press, 526p. Desert Research Institute (DRI), 2014, Monthly Climate Summary - Western Regional Climate Center data for Calexico, Period of Record 1/1/1904 to 6/30/2010. Dohrenwend, J.C., and Smith, R.S.U., 1991, Quaternary Geology and Tectonics of the Salton Trough: in Morrison, R.B. (Editor), Quaternary Nonglacial Geology: Conterminous U.S.: Geological Society of America, Boulder, Colorado, The Geology of North America, v. K-2, p. 334-337. Havholm, K.G., and Kocurek, G., 1988, A Preliminary Study of the Dynamics of a Modern Draa, Algodones, Southeastern California, USA: Sedimentology, v. 35, p. 649-669. Jennings, C.W., 1967, Geologic Map of California, Salton Sea Sheet: California Division of Mines and Geology, scale 1:250,000. Kocurek, G., and Nielson, J., 1986, Conditions Favorable for the Formation of Warm-Climate Aeolian Sand Sheets: Sedimentology, v.33, p. 795-816. Lancaster, N., and Tchakerian, V.P., 2003, Late Quaternary Eolian Dynamic, Mojave Desert, California: in Enzel, Y., Wells, S.G., and Lancaster, N. (Editors), Paleoenvironments and Paleohydrology of the Mojave and Southern Great Basin Deserts: Geological Society of America, Special Paper 368, p. 231-249. Lancaster, N., 1988, Development of Linear Dunes in the Southwestern Kalahari, Southern Africa: Journal of Arid Environments, 14(3), 233-244. Lancaster, J. T., and Bedrossian, T. L., 2014, Memorandum Report Dated August 2014, Prepared for Mr. Serge Glushkoff, CDFW, Titled, Eolian System Mapping for the Desert Renewable Energy Conservation Plan, California Geological Survey,

54p., 4 plates (multiple map scales). Loeltz, O.J., Irelan, B., Robinson, J.H., and Olmstead, F.H., 1975, Geohydrologic Reconnaissance of the Imperial Valley, California: U.S. Geological Survey, Professional Paper 486-K. McKee, E.D., (Editor), 1979, A Study of Global Sand Seas: U.S. Geological Survey, Professional Paper 1052, 429 p. Muhs, D.R., Bush, C.A., Cowherd, S.D., and Mahan, S., 1995, Geomorphic and Geochemical Evidence for the Source of Sand in the Algodones Dunes, Colorado Desert, Southeastern California: in Tchakerian, V.P., Desert Aeolian Processes: Chapman & Hall, London, p. 57-74. Muhs, D.R., Reynolds, R.L., Been, J., and Skipp, G., 2003, Eolian Sand Transport Pathways in the Southwestern United States: Importance of the Colorado River and Local Sources: Quaternary International, v. 104, p. 3-18. Nielson, J., and Kocurek, G., 1986, Climbing Zibars of the Algodones: Sedimentary Geology, v. 48, p. 1-15. Norris, R.M., and Norris, K.S., 1961, Algodones Dunes of Southeastern California: Geological Society of America Bulletin, v. 72, p. 605-620. Norris, R. M., and Webb, R. W., 1990, Geology of California: Wiley. OlmsteadF.H., Loeltz, O.J., and Irelan, B., 1973, Geohydrology of the Yuma Area, Arizona and California: U.S. Geological Survey, Professional Paper 486-H.

Sharp, R.P., 1979, Intradune Flats of the Algodones Chain, Imperial Valley, California: Geological Society of America Bulletin, v. 90, p. 908-916. Smith, R.S.U., 1982, Sand Dunes in the North American Deserts: in Bender, G.L. (Editor) Reference Handbook on the Deserts of North America: Greenwood Press, Westport, Connecticut, p. 481-524. Soil Survey Staff, 1903, Soil Survey of Imperial Area, Imperial County, California, p1219-1248. Soil Survey Staff, 1981, Soil Survey of Imperial County, California, Imperial Valley Area, 112p. Sweet, M.L., Nielson, J., Havholm, K., and Farrelley, J., 1988, Algodones Dune Field of Southeastern California: Case History of a Migrating Modern Dune Field: Sedimentology, v. 35, p. 939-952. Van Andel, T.H., 1964, Recent Marine Sediments of the Gulf of California: in, Van Andel, T.J., Shor, G.G. (Eds.), Marine Geology of the Gulf of California: A Symposium. AAPG Mem., vol. 3, pp. 216–310.

Van de Kamp, P.C., 1973, Holocene Continental Sedimentation in the Salton Basin, California: A Reconnaissance: Geological Society of America Bulletin, v. 84, p. 827-848. Waters, M. R., 1983, Late Holocene Lacustrine chronology and Archaeology of Ancient Lake Cahuilla, California. Quaternary Research, 19(3), 373-387. Zimbleman, J.R., Williams, S.H., and Tchakerian, V.P., 1995, Sand Transport Paths in the Mojave Desert, Southwestern United States: in Tchakerian, V.P. (editor), Desert Aeolian Processes: Chapman and Hall, London, p. 101-129. Imagery Used

Google Earth, 2005, Digital Globe Imagery, accessed via Google Earth (v. 7.1.2.2041) March - April, 2014 U.S. Department of Agriculture, Farm Service Agency-Aerial Photography Field Office, National Agriculture Imagery Program (NAIP), 2005, 1-meter resolution. U.S. Department of Agriculture, Farm Service Agency-Aerial Photography Field Office, National Agriculture Imagery Program (NAIP), 2009, 1-meter resolution. U.S. Department of Agriculture, 2012, accessed via Google Earth (v. 7.1.2.2041) March - April, 2014.

PLATE 4

Predominant transport direction. Inferred from geomorphic interpretation of dune crest migration, sand sheets and streaks, and ventifacts. Active sand transport